

4D-Var for Dummies

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Centre for Australian Weather and Climate Research
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#1

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*The Fun and Easy Way[™]
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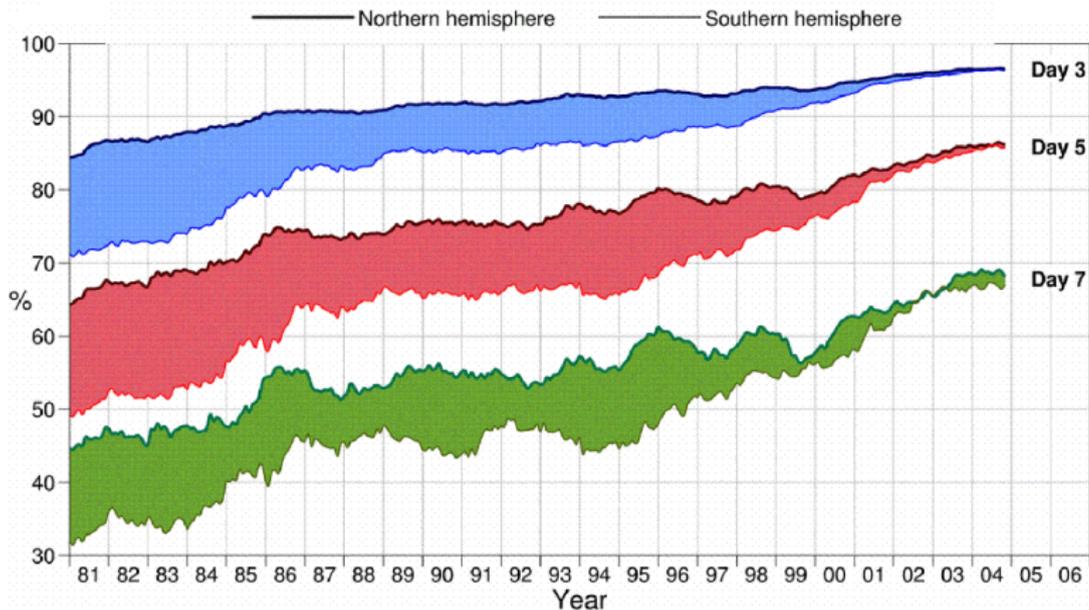
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- ▶ 1854: Meteorological Dept of the British Board of Trade created
- ▶ “...in a few years we might know in this metropolis the condition of the weather 24 hours beforehand.” (M.J.Ball MP, House of Commons, 30 June 1854.)
- ▶ Response from House: “Laughter”

Anomaly correlation of ECMWF 500hPa height forecasts



Why Data Assimilation is Important

- ▶ Numerical Weather Prediction (NWP) is (largely) an initial value problem.
 - ▶ Has contributed to enormous forecast improvements
 - ▶ Extracts the maximum value from expensive observations
- ▶ Accurate analyses are necessary for getting the most from field programs.
- ▶ Reanalyses of past data using modern methods are an essential resource for climate research.

Best Linear Unbiased Estimate (BLUE)

- ▶ Observations y_1 and y_2 of a true state x_t :

$$y_1 = x_t + \epsilon_1$$

$$y_2 = x_t + \epsilon_2$$

- ▶ The statistical properties of the errors are known:

$$\langle \epsilon_1 \rangle = 0$$

$$\langle \epsilon_1^2 \rangle = \sigma_1^2$$

$$\langle \epsilon_1 \epsilon_2 \rangle = 0$$

$$\langle \epsilon_2 \rangle = 0$$

$$\langle \epsilon_2^2 \rangle = \sigma_2^2$$

- ▶ **Estimate** x_a of x_t as a **linear** combination of the observations such that $\langle x_a \rangle = x_t$ (**unbiased**) and $\sigma_a^2 = \langle (x_a - x_t)^2 \rangle$ is minimised (**best**).
- ▶ Then

$$x_t = \frac{\sigma_2^2 y_1 + \sigma_1^2 y_2}{\sigma_1^2 + \sigma_2^2}$$

Best Linear Unbiased Estimate (cont'd)

- ▶ Same estimate found by minimising

$$J(x_a) = \frac{(x_a - y_1)^2}{\sigma_1^2} + \frac{(x_a - y_2)^2}{\sigma_2^2}$$

- ▶ Minimising J is the same as maximising $\exp(-J/2)$
- ▶ Hence for Gaussian errors the BLUE is the maximum likelihood (or optimal) estimate.
- ▶ For many pieces of data $\mathbf{y} = (y_1, y_2, \dots, y_n)^T$,

$$J(x_a) = (x_a - \mathbf{y})^T \mathbf{P}^{-1} (x_a - \mathbf{y})$$

where \mathbf{P} is the error covariance matrix of \mathbf{y} .

Assimilation: The Big BLUE

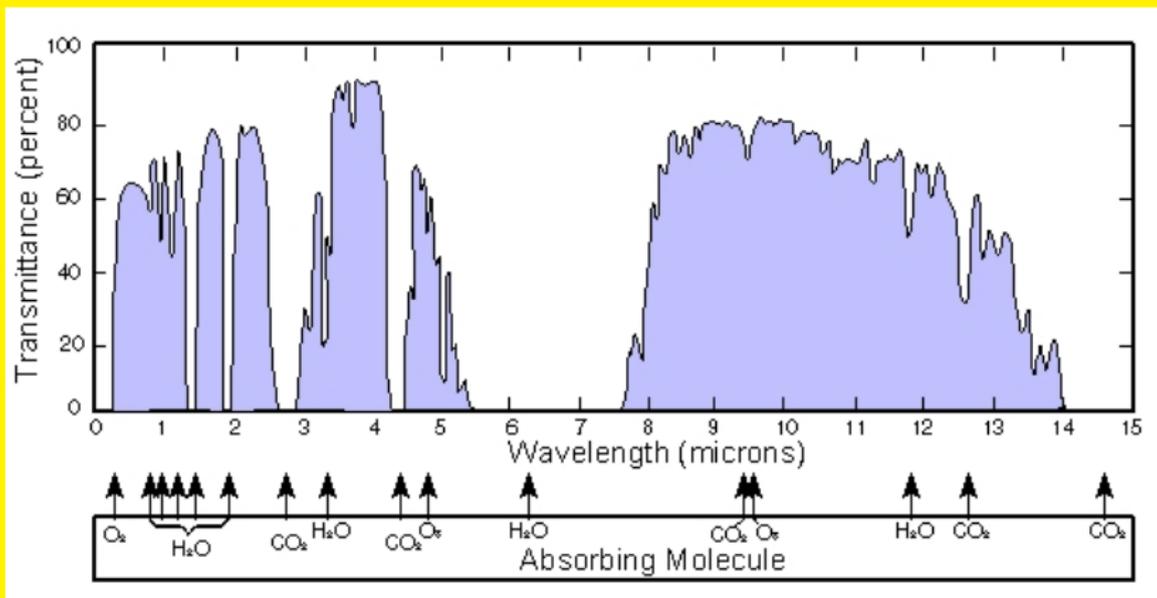


Assimilation combines a short-term numerical forecast with some observations:

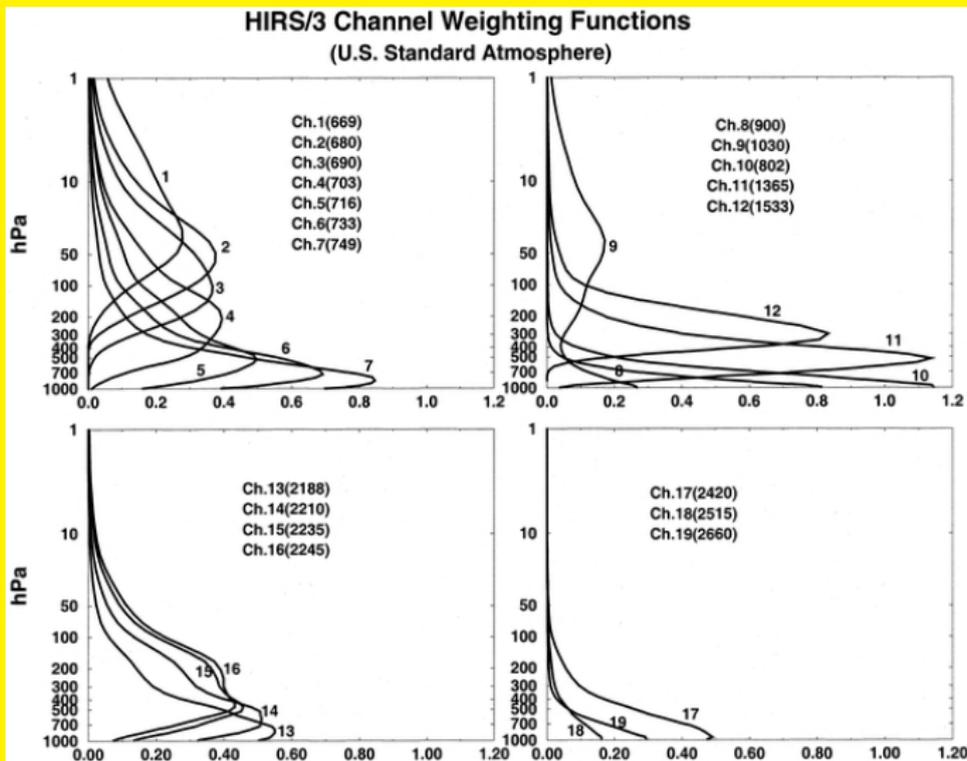
$$J(\mathbf{x}_a) = (\mathbf{x}_a - \mathbf{x}_f)^T \mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_f) + (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})$$

- ▶ \mathbf{x}_a is the analysis
- ▶ \mathbf{x}_f the short-term forecast
- ▶ \mathbf{y} are the observations
- ▶ \mathcal{H} produces the analysis estimate of the observed values
- ▶ \mathbf{R} is the observation error covariance
- ▶ \mathbf{B} is the forecast error covariance

The Atmospheric Infrared Transmission Spectrum



HIRS (High resolution InfraRed Sounder) Channel Weights



Finding the minimum of J

$$J(\mathbf{x}_a) = (\mathbf{x}_a - \mathbf{x}_f)^T \mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_f) + (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})$$

Solve directly $\nabla J = 0$.

- ▶ Have to manipulate big matrices
- ▶ Nonlinear \mathcal{H} is very difficult (satellite radiances)

Iterative minimisation (a.k.a. **variational assimilation**).

- ▶ Finds full 3-D structure of the atmosphere (**3D-Var**)
- ▶ Other observations and background helps constrain the poorly-conditioned and underdetermined inversion of the satellite radiances

Minimising J

$$J(\mathbf{x}_a) = (\mathbf{x}_a - \mathbf{x}_f)^T \mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_f) + (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})$$

To minimise J , we need the gradient:

$$\nabla J(\mathbf{x}_a) = 2\mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_f) + 2\mathbf{H}^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})$$

$\mathbf{H} = \left[\frac{\partial \mathcal{H}_i}{\partial \mathbf{x}_{a,j}} \right]$ is the Jacobian of \mathcal{H} (a.k.a. the **tangent linear**)

\mathbf{H}^T is the **adjoint** of \mathcal{H}



The Importance of \mathbf{B}

$$J(\mathbf{x}_a) = (\mathbf{x}_a - \mathbf{x}_f)^T \mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_f) + (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})$$

\mathbf{B} is important:

- ▶ Conditioning and speed of convergence
- ▶ Getting the statistics right
- ▶ Describing atmospheric balance
- ▶ Spatial scale of analysis

\mathbf{B} in the model variables fails miserably:

- ▶ Rank deficient
- ▶ Too large to store, let alone operate on



B the best you can **B**



Representing **B** typically involves:

- ▶ *Transform to less-correlated variables.*
 - ▶ $(u, v) \implies (\psi, \chi)$
 - ▶ $u = -\partial\psi/\partial y + \partial\chi/\partial x$, $v = \partial\psi/\partial x + \partial\chi/\partial y$
 - ▶ Replace mass field by unbalanced mass:
$$\phi_{\text{unbal}} = \phi - \phi_{\text{bal}}(\psi)$$
- ▶ *Transform to spectral space.*
- ▶ *Rescale.*

These make **B** diagonal \implies good conditioning and computational efficiency.

- ▶ *Truncate the small scales.* Forecast error spectrum is red, with little power at small scales. So truncate **B**.

Incremental Formulation

Replace

$$J(\mathbf{x}_a) = (\mathbf{x}_a - \mathbf{x}_f)^T \mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_f) + (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})$$

by

$$J(\delta\mathbf{x}) = \delta\mathbf{x}^T \mathbf{B}^{-1} \delta\mathbf{x} + (\mathcal{H}(\mathbf{x}_f) + \mathbf{H}\delta\mathbf{x} - \mathbf{y})^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_f) + \mathbf{H}\delta\mathbf{x} - \mathbf{y})$$

where $\delta\mathbf{x} = \mathbf{x}_a - \mathbf{x}_f$ and \mathbf{H} is the Jacobian of \mathcal{H} (**tangent linear**).

- ▶ $\mathcal{H}(\mathbf{x}_a)$ becomes $\mathcal{H}(\mathbf{x}_f) + \mathbf{H}\delta\mathbf{x}$
- ▶ Computational efficiency since $\delta\mathbf{x}$ now at reduced resolution of \mathbf{B} , \mathbf{H} maybe cheaper to compute than \mathcal{H} , true quadratic form.



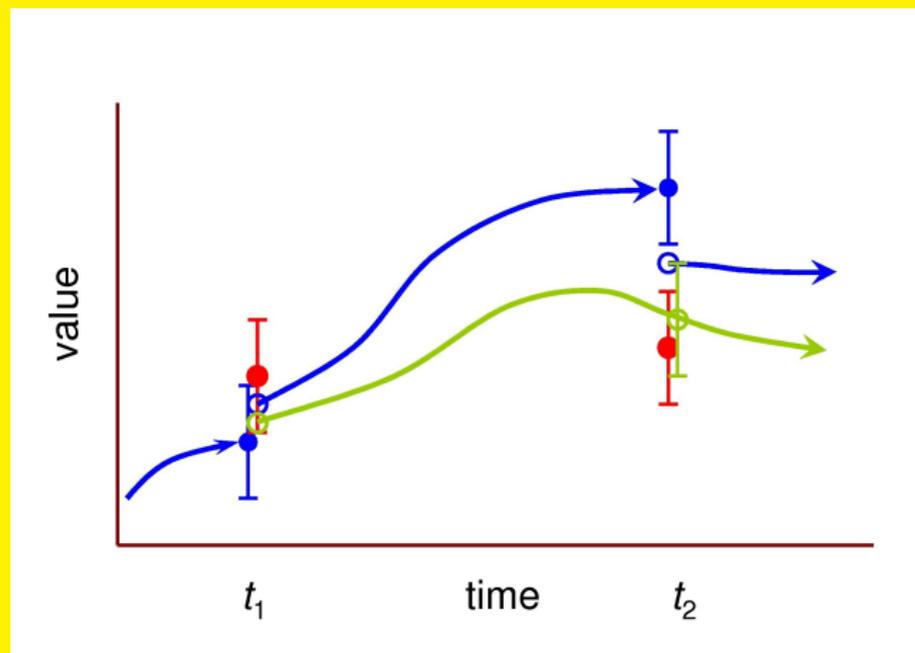
A Matter of Time

So far all data assumed to be at the analysis time.

- ▶ Assimilate e.g. four times a day.
- ▶ All data in 6-hour window assumed to occur at the middle of that window.
- ▶ Introduces some errors \implies weather systems move and develop!
- ▶ Reduce errors by assimilating more frequently, but that has its own problems.

A better way is to introduce the time dimension into the assimilation, 4-dimensional variational assimilation (**4D-Var**).

Observations at two times



Red: Observations. Blue: 3D-Var. Green: 4D-Var.

4D-Var

Add a term for the later time:

$$J(\mathbf{x}_a) = \dots + (\mathcal{H}_2(\mathcal{M}(\mathbf{x}_a)) - \mathbf{y}_2)^T \mathbf{R}_2^{-1} (\mathcal{H}_2(\mathcal{M}(\mathbf{x}_a)) - \mathbf{y}_2)$$

- ▶ \mathcal{M} is the model forecast from t_1 to t_2
- ▶ Subscripts 2 refer to the time t_2 .

The gradient becomes

$$\nabla J(\mathbf{x}_a) = \dots + 2\mathbf{M}^T \mathbf{H}_2^T \mathbf{R}_2^{-1} (\mathcal{H}_2(\mathcal{M}(\mathbf{x}_a)) - \mathbf{y}_2)$$

- ▶ \mathbf{H}_2^T is the adjoint of the Jacobian of \mathcal{H} , takes information about the observation-analysis misfit from radiance space to analysis space
- ▶ \mathbf{M}^T is the adjoint of the Jacobian of \mathcal{M} and propagates this gradient information back in time from t_2 to t_1 .

4D-Var

Minimising

$$\begin{aligned} J(\mathbf{x}_a) &= (\mathbf{x}_a - \mathbf{x}_f)^T \mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_f) \\ &\quad + (\mathcal{H}(\mathbf{x}_a) - \mathbf{y})^T \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}_a) - \mathbf{y}) \\ &\quad + (\mathcal{H}_2(\mathcal{M}(\mathbf{x}_a)) - \mathbf{y}_2)^T \mathbf{R}_2^{-1} (\mathcal{H}_2(\mathcal{M}(\mathbf{x}_a)) - \mathbf{y}_2) \end{aligned}$$

gives an analysis \mathbf{x}_a at time t_1 that

- ▶ is close to the background \mathbf{x}_f at t_1
- ▶ is close to the observations \mathbf{y} at t_1
- ▶ initialises a (linearised) forecast that is close to the observations \mathbf{y}_2 at time t_2

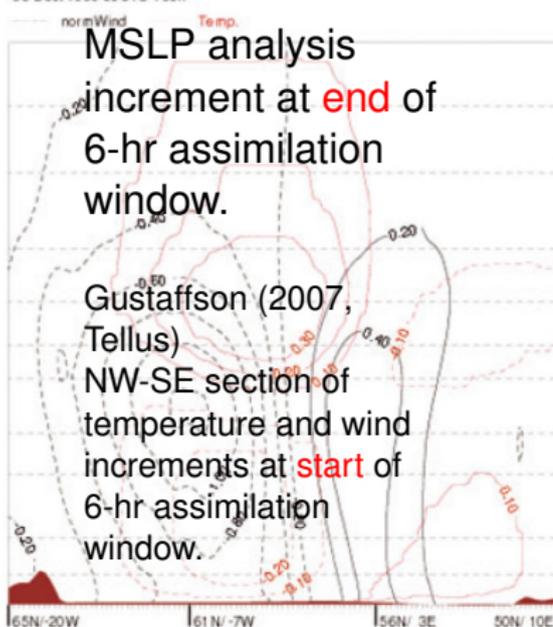
Adding additional time levels is straightforward, as is the incremental formulation (**exercise**).

4D-Var analysis of a single pressure observation

One pressure observation at centre of low, 5hPa below background, at end of 6-hr assimilation window.



03 Dec. 1999 06UTC +00h



In practice ...

This is not a small problem!

- ▶ Atmospheric model has $O(10^6 \text{ to } 10^7)$ variables
- ▶ Millions of observations per day
- ▶ Limited time available under operational constraints

The model has several hundred thousand lines of code, 4D-Var requires

- ▶ operations by the Jacobian of the model
- ▶ operations by the adjoint of the Jacobian

Good results require accurately estimating the necessary statistics (**R** and **B**) and careful quality control of the observations.

Extensions

Multiple “Outer Loops”

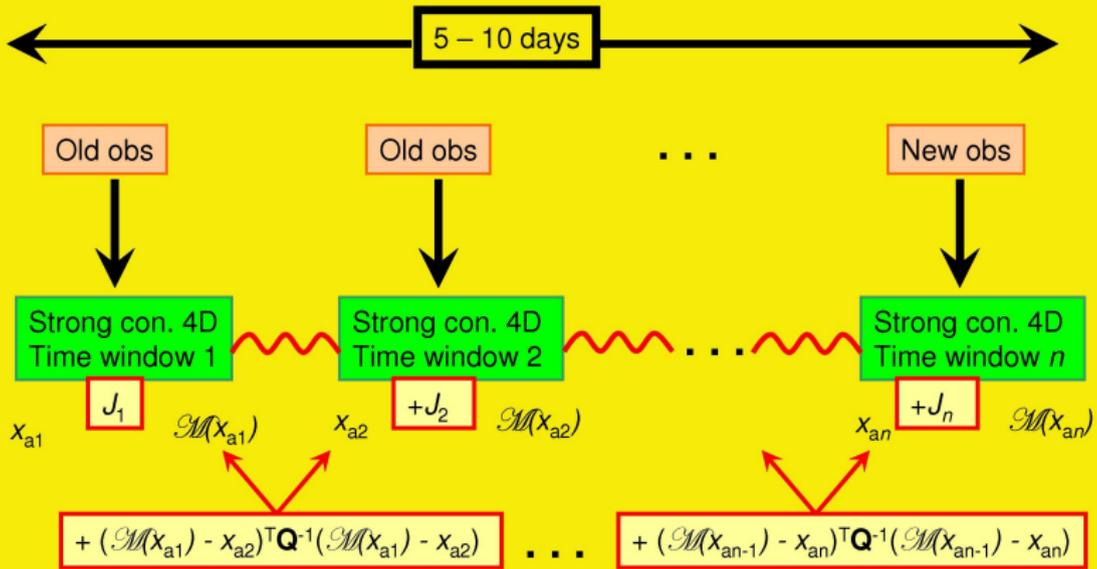
- ▶ *Problem:* Accuracy is limited by the linearisations of \mathcal{H} and \mathcal{M} .
- ▶ *Solution:* Update the nonlinear forecast (outer loop) several times during the minimisation of the $J(\delta\mathbf{x})$ (inner loop).

Multi-incremental 4D-Var

- ▶ *Problem:* Balancing speed of convergence against need to resolve small scales.
- ▶ *Solution:* Begin minimising with $\delta\mathbf{x}$ at low resolution, and increase resolution after each iteration of the outer loop.

Extensions: Weak Constraint 4D-Var

- ▶ Doesn't assume that the model is perfect
- ▶ Allows a longer window.



Summary

Var is better than direct solution (a.k.a. Optimum Interpolation) because:

- ▶ Can handle lots of observations
- ▶ Can better cope with nonlinear observation operator \mathcal{H}
- ▶ Solves for the whole domain at once

4D-Var is better than 3D-Var because:

- ▶ Uses observations at the correct time
- ▶ Calculates analysis at the correct time
- ▶ Implicitly generates flow-dependent **B**

Special issues of QJRMS and JMetSocJapan from WMO DA workshops.

Kepernt, J.D., 2007: Maths at work in meteorology. *Gazette of the Australian Mathematical Society*, **34**, 150 – 155. Available from

<http://www.austms.org.au/Publ/Gazette/>